

1 Lab Methods

Hydrogeologic Parameters: Grain Size

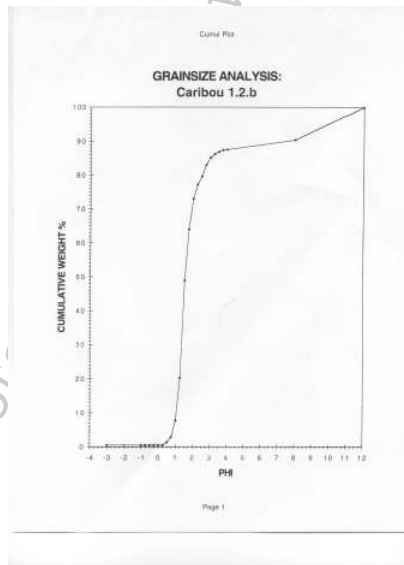
- Grain Size Analysis, Sandy material
- Hazen's equation: $K = C(d_{10}^2)$,
 - C constant ranging from 100 to 150 $\text{sec} \cdot \text{cm}^{-1}$
 - d_{10} is screen size (10% passing)
- assumes uniformly graded (poorly sorted) sands ($\frac{d_{60}}{d_{10}} < 5$)

Hydrogeologic Parameters: Grain Size

- Kozeny-Carmen equation
- Several variation with different assumptions
- For spherical particles with radius= r

$$K = \frac{\rho g}{\mu} \frac{n^3}{(1-n)^2} \frac{r^2}{45}$$
$$K = \frac{\rho g}{\mu} \frac{n^3}{(1-n)^2} \frac{d_{50}^2}{180}$$

Grain Size Curve

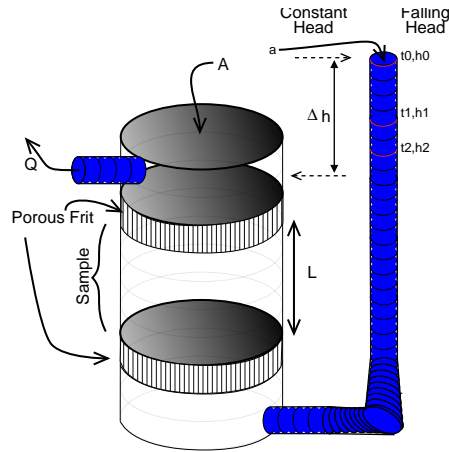


Hydrogeologic Parameters: Permeameters

- Constant head permeameter
- Measure hydraulic conductivity of permeable sediments
- based on the simple rearrangement of Darcy's Law

$$K = \frac{QL}{AH}$$

Hydrogeologic Parameters: Permeameters



Hydrogeologic Parameters: Permeameters

- Falling head permeameter
- for lower permeability sediments
- measure rate that water falls in a manometer
- head gradient discharge through sample change with time
- set flux based on Darcy's Law equal to the volume change in the manometer

$$KA \frac{h(t) - H_0}{L} = a \frac{dh}{dt}$$

$$K = \frac{aL}{A(t_2 - t_1)} \ln \left[\frac{h_2}{h_1} \right]$$

2 Slug Testing

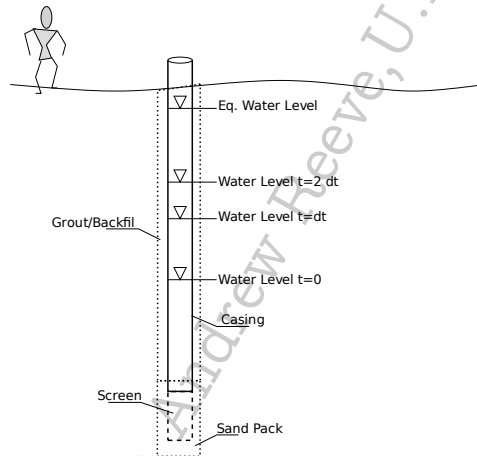
Piezometer Testing

- Measure pre-test water level in well
- Introduce (or remove) 'slug', rapidly change water level
- Measure water level recovery with time
- Rate of recovery related to hydraulic conductivity

Piezometer Testing

- Information needed:
 - Equilibrium water level
 - Water levels and times during recovery
 - Volume of 'slug'
 - Well geometry
 - * screen length
 - * well casing radius
 - * sand pack or screen radius
 - * screen length
 - * distance between screen and water table
 - location of confining layers

Piezometer Testing



Hvorslev Method

- Simplest method for interpreting piezometer tests
- Based on water balance between flux across screen and volume changing in casing

$$q = \pi \cdot r_c^2 \frac{dh}{dt}$$

$$q = -F \cdot K (h_{eq} - h_i)$$

$$F = \frac{A}{\Delta r}$$

$$T_0 = \frac{\pi \cdot r_c^2}{F \cdot K}$$

$$\pi \cdot r_c^2 \frac{dh}{dt} = -F \cdot K (h_{eq} - h_i)$$

$$-\int_{t_0}^{t_i} \frac{1}{T_0} dt = \int_{t_0}^{t_i} \frac{1}{h_{eq} - h_i} dh$$

$$-\frac{\Delta t}{T_0} = \ln \left(\frac{h_{eq} - h_i}{h_{eq} - h_0} \right)$$

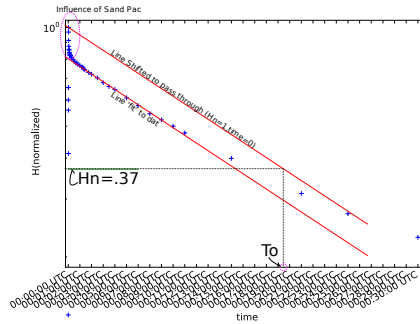


Figure 1: Example of graphical analysis by the Hvorslev method. The linear fit to the data can be used to determine the lag time (T_0) or used to find two points along the best fit line. The inverse of the lag time should equal $\ln\left(\frac{H_1}{H_2}\right) \frac{1}{\Delta t}$. Note semilog graph.

Hvorslev Example

Cooper et al. Method (WRR,1967)

- Cooper et al. (1967) solved radial flow equation for instantaneous source.
 - infinite extent
 - significant storage in well and aquifer
 - fully penetrating well

$$\frac{H}{H_0} = \frac{8\alpha}{\pi^2} \int_0^{\infty} \frac{\exp\left(\frac{-\beta u^2}{\alpha}\right)}{u \cdot \Delta u}$$

•

$$\Delta u = (u \cdot J_0(u) - 2 \cdot \alpha \cdot J_1(u))^2 + (u \cdot Y_0(u) - 2 \cdot \alpha \cdot Y_1(u))^2$$

$$\alpha = \frac{r_s^2 S_s b}{r_c^2}$$

$$\beta = \frac{Kbt}{r_c^2}$$

Cooper et al. Method

- Cooper et al (1967) presented type curves for different values of alpha and beta
- plot normalized head (arithmetic scale, y-axis) vs. time (log scale, x-axis)
- overlay data plot on type curves, slide along x-axis until data matches a type curve
- pick time and note corresponding beta value, note alpha value of matched type curve
- calculate T and S
- Note that S lacks accuracy

Cooper et al. Example

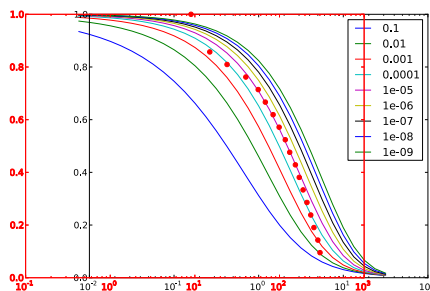
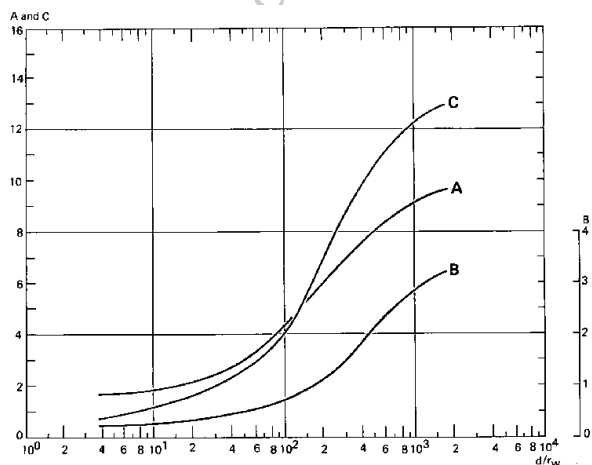


Figure 2: Type curves (black axes, β vs H_n) overlain with data (red axes, time (sec) vs H_n). Data is obviously of poor quality, and illustrates the difficulty in picking a correct type curve. Here, $\alpha = 10^{-5}$ and one possible match point on the x-axis is (time=1, $\beta \approx 0.018$).

Bouwer-Rice Method (WRR,1976)

- Slug testing method developed for unconfined aquifers
- Similar to Hvorselv method (plotting)
 - plot $\log(\text{normalized head})$ vs. time
 - pick value from straight portion of curve and get t, h_t
 - use ratio of l/r_s and type curves to calculate A, B, and C.
 - use appropriate eqn to calculate K

Bouwer-Rice Method



Bouwer-Rice Method

$$r_s^* = r_s \sqrt{\frac{K_v}{K_h}}$$

$$A = 1.472 + 3.537 \cdot 10^{-2} \frac{b}{r_s^*} - 8.148 \cdot 10^{-5} \frac{b^2}{r_s^{*2}} + 1.028 \cdot 10^{-7} \frac{b^3}{r_s^{*3}} - 6.484 \cdot 10^{-11} \frac{b^4}{r_s^{*4}} + 1.573 \cdot 10^{-14} \frac{b^5}{r_s^{*5}}$$

$$B = 0.2372 + 5.151 \cdot 10^{-3} \frac{b}{r_s^*} - 2.682 \cdot 10^{-6} \frac{b^2}{r_s^{*2}} - 3.491 \cdot 10^{-10} \frac{b^3}{r_s^{*3}} + 4.738 \cdot 10^{-13} \frac{b^4}{r_s^{*4}}$$

$$C = 0.7920 + 3.993 \cdot 10^{-2} \frac{b}{r_s} - 5.7443 \cdot 10^{-5} \frac{b^2}{r_s^2} + 3.858 \cdot 10^{-8} \frac{b^3}{r_s^3} - 9.659 \cdot 10^{-12} \frac{b^4}{r_s^4}$$

Bouwer-Rice Method

- Calculate effective radius

– for partially penetrating wells (H_{br} = distance from bottom of screen to water table.):

$$\ln \frac{R_e}{r_s} = \left[\frac{1.1}{\ln\left(\frac{H_{br}}{r_s}\right)} + \frac{A + B \ln\left[\frac{b - H_{br}}{r_s}\right]}{\frac{1}{r_s}} \right]^{-1}$$

- if $\ln\left(\frac{b - H_{br}}{r_s}\right)$ is greater than 6.0, then 6.0 should be used
- for fully penetrating wells:

$$\ln \frac{R_e}{r_s} = \left[\frac{1.1}{\ln\left(\frac{H_{br}}{r_s}\right)} + \frac{C}{\frac{1}{r_s}} \right]^{-1}$$

- calculate K

$$K = \frac{r_c^2 \ln \frac{R_e}{r_s}}{2l} \frac{1}{t} \ln \frac{h_0}{h_t}$$

$$K = \frac{r_c^2 \ln \frac{R_e}{r_s}}{2lT_0}$$

Notes from Andrew Reeve, U. Maine